

Inter-model differences in the representation of the AMOC forcing of the NAO

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Thermohaline Circulation (THC)





Atlantic Meridional Overturning Circulation (AMOC)

RAPID array monitoring AMOC strength at 26N since 2004







MSLP correlation with Nov-Mar NAO index 1979-2019

North Atlantic Oscillation (NAO)

Dominant mode of variability over the North Atlantic

Observations based on ERA5 reanalysis

15 HighResMIP models

100-year control simulations

1950's conditions

RAPID-ERA5 2004-2022

ERA5 1941-2023

L_{eff} (km)	L_{nom} (km)	O_{res} (degrees)
\geq 625	250	1
364	100	1/4
185	50	1/4
${\leq}185$	50	1/12
571	100	1
182	25	$1/_{4}$
253	50	1
\geq 185	50	1/4
185	25	1/4
351	100	1
238	50	1/4
364	100	1
256	50	1/4
\geq 625	250	1
313	50	1/4
	$\begin{array}{r c c c c c c c c c c c c c c c c c c c$	$\begin{array}{ c c c c c } L_{eff} (\mathbf{km}) & L_{nom} (\mathbf{km}) \\ \hline \geq 625 & 250 \\ \hline 364 & 100 \\ 185 & 50 \\ \leq 185 & 50 \\ \hline 571 & 100 \\ 182 & 25 \\ 253 & 50 \\ \geq 185 & 50 \\ 185 & 25 \\ 351 & 100 \\ 238 & 50 \\ \hline 364 & 100 \\ 256 & 50 \\ \geq 625 & 250 \\ \hline 313 & 50 \\ \end{array}$



- Winter season (DJF)
- Nonlinear detrended
- AMOC index: strength at 26N in Sv
- Different lags using 5 year running means
 - Negative lag: atmosphere is before ocean
 - Positive lag: atmosphere is lagging ocean





NAO MSLP response lag correlated with AMOC



Large spread among PRIMAVERA models especially at positive lags Separation between models according to their response at lag +5

NAO MSLP response lag correlated with AMOC







RAPID – ERA5



Smith et al. 2020 "North Atlantic climate far more predictable than models imply"



Regression MSLP on AMOC

Neg. NAO models Model mean

Pos. NAO models Model mean

regression

RAPID-ERA5 2004-2022



Lag +5







Lag 0



regression psl NAO-Modelmean lag+0 rm5



regression MOC msl RAPID ERA5 lag+0 rm5



Lag -3



regression psl NAO-Modelmean lag-3 rm5



regression MOC msl RAPID ERA5 lag-3 rm5





At lag 0 the AMOC response is dominated by the Ekman transport

Polo et al. 2014. Ocean model forced with ERA-40 winds





regression tos NAO+Modelmean lag+5 rm5





SST averaged over SPG box regressed on AMOC

Neg. NAO models



Lag in years

Pos. NAO models





-0.25

0.50

-0.75

-1.00 -

-20

-15

-10

-5

RAPID-ERA5



Lag in years

Lag in years

Time Lags

10

15

20



Regression SST on AMOC

Neg. NAO models Regression turbulent heat flux on AMOC (positive upward)









Lag -3

12



Regression turb.

heatflux on AMOC

Regression SST on AMOC

Rapid – ERA5



Lag -3



ERA5 1981-2023 5 yr running mean



Regression SST SPG box on SST, THF and MSLP





How is the atmosphere forced by the lagged SST response?





AMOC forcing and evolution of SSTs

Pos. NAO models

Mixed Layer Depth

Neg. NAO models



Regression with AMOC at Lag -3





Bias







regression tos NAO-Modelmean lag-3 rm5



regression MOC sst RAPID ERA5 lag-3 rm5







regression tos NAO-Modelmean lag+0 rm5





Lag + 3





Regression



Pos. NAO





What determines the difference between the neg. NAO and pos. NAO models?

- Bias
- Resolution



Bias

Sea-ice conc.



SST



T2m



Neg. NAO models

Pos. NAO models larger bias than Neg. NAO models







Pos. NAO models



SPG bias \longleftrightarrow NAO response at lag +5



Neg. NAO models Other models



Kim et al. 2023

γ: sensitivity of deepwater formation tobuoyancy flux

 β : sensitivity of AMOC response to NAO forcing



weak bias

Neg. NAO models

Pos. NAO models



Regression on AMOC at lag -3



regression tas NAO-Modelmean lag-3 rm5



Z500

regression tas NAO+Modelmean lag-3 rm5

At lag-3 T2m response seems to affect MSLP response

> Cold – High Warm – Low mechanism





Impact of sea-ice response

- 0.6

0.0

-0.3

-0.6

Effective resolution atmosphere models

(Klaver et. al 2020)



${\rm High \ Ocean} \ (0.25^{\circ})$					Low Ocean (1°)						
$\label{eq:VHigh Res} {\bf V} \ {\bf High} \ {\bf Res} \ (\leq 185 \ {\bf km}) \ \ {\bf High} \ {\bf R}$			High Res	ligh Res (185-256 km)		Low Res (256-364 km)		V Low Res (\geq 364 km)			
model	group	SPG SST bias	model	group	SP SST bias	model	group	SPG SST bias	model	group	SPG SST bias
ECMWF HR	Neg. NAO	-1.3	ECMWF MR	Pos. NAO	-1.0	CNRM HR	Pos. NAO	-5.9	HadGEM LL	Neg. NAO	-3.3
HadGEM HH	Neg. NAO	+0.6	MPI XR	Pos. NAO	-3.2	HadGEM MM	NAO0	-0.3	CNRM LR	Neg. NAO	+1.2
HadGEM HM	Neg. NAO	+0.1	ECMWF LR	Pos. NAO	-5.9	MPI HR	NAO0	-2.2			
CMCC-VHR	Neg. NAO	+1.1	EC-Earth HR	Pos. NAO	-1.8		Pos. NAO	-2.4			

SPG SST bias is averaged over SPG box in °C

Neg. NAO response for very high and very low-resolution models

Very low-resolution models well-tuned?

DWF for a few models at lag -3

HadGEM HH



HadGEM HM



HadGEM MM



HadGEM LL

regres AMAX1 mlotst HadGEM LL lag-3 rm5 nonlin_detr

CNRM LR



EC-Earth HR





NAO – AMOC feedback



Scatter plot DWF lag -3 with NAO response at lag +5



NAO - AMOC feedback is non-linear



All AMOC

Neg. AMOC

Pos. AMOC

Time Lags



Conclusions



- Different behaviour of PRIMAVERA models with respect to AMOC-NAO interaction
- SST bias in the SPG appears to be the main cause for the different behaviour. Resolution is important to reduce the bias. Tuning can compensate errors.
- Models with small bias simulate a switch from positive to negative NAO at positive lags of the AMOC
- RAPID ERA5 and ERA5 support the forcing of a negative NAO by the AMOC at positive lags

Discussion

For reliable decadal predictions:

- Crucial to reduce the SST-bias in the SPG and the sea-ice bias in the GIN sea
- Use flux correction to reduce the bias?

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